

Carbon flux through the microbial community in a temperate sea during summer: role of bacterial metabolism

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ABSTRACT: The carbon flux through bacterioplankton under oligotrophic conditions was studied during summer in the southern Bay of Biscay. An episode of coastal upwelling during our study allowed us to look at how bacterial metabolism responded to short-term environmental changes. Gross primary production (GPP) was significantly higher after the upwelling pulse; however, microbial respiration (R) exceeded the organic carbon fixation by primary producers ($R > \text{GPP}$), resulting in net heterotrophic balance during summer (net community production [NCP] < 0). Heterotrophic bacteria were the major contributors to bulk community respiration, accounting for 70% of the total respiration. The variability in both bacterial respiration and plankton respiration was largely attributable to differences in the dissolved organic carbon (DOC) concentration. The ratio of bacterial biomass to total phytoplankton biomass decreased after the upwelling pulse, whereas the bacterial production to primary production ratio increased as the productivity of the system increased. Bacterial production was always significantly lower than bacterial respiration, giving rise to a bacterial growth efficiency (BGE) lower than 10%. BGE responded to changes in phytoplankton production, increasing as the system became more productive. Short-term variations in the structure of the water column may have an important impact on bacterial processes within this oligotrophic system.

KEY WORDS: Bacterial respiration · Bacterial production · Bacterial growth efficiency · Gross primary production · Upwelling pulse · Oligotrophic period

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INTRODUCTION

Bacterioplankton are the major consumers of organic matter, and represent a significant biomass component in aquatic planktonic ecosystems. According to a recent review (Cotner & Biddanda 2002), the relative contribution of bacterioplankton to organic-matter cycling and energy is greatest in oligotrophic ecosystems. For this reason, knowing the factors that regulate the marine bacterial metabolism is ecologically and biogeochemically important in order to understand the cycling of materials within the oligotrophic regions.

Planktonic bacteria are responsible for the bulk of microbial respiration (Cole et al. 1988). However, of all the aspects of bacterial metabolism, bacterial net production (BNP) has traditionally received more atten-

tion than bacterial respiration (BR), probably due to the methodological difficulties involved in the study of the latter (Kirchman 1997). Bacterial growth efficiency (BGE) is the amount of new bacterial biomass produced per unit of organic C substrate assimilated, and is related to the quality of the available dissolved organic matter (DOM), the nutrient availability and the energy demand of the system (del Giorgio & Cole 1998). There is a large range of BGE variation in natural planktonic bacteria, from less than 5% to more than 60%. However, the lack of simultaneous BNP and respiration measurements, together with the fact that respiration rates are typically derived from biomass and production measurements (Jahnke & Craven 1995), makes it difficult to assess the organic carbon flow through bacterioplankton and therefore hampers

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our knowledge of the functioning of planktonic metabolism. Most studies providing estimates of BGE have been carried out in productive areas while data in oligotrophic waters are scarce (see references in del Giorgio & Cole 1998). On the other hand, the lack of information on the temporal variability of bacterial metabolism and BGE may be another important limitation to our understanding of organic carbon fluxes through the microbial community in the ocean. The temporal variability in bacterial metabolism is often not considered, and a single data-point is considered to be representative of extended time-periods, the results of which can only be interpreted as representing constant conditions of the systems studied.

The Bay of Biscay displays thermal stratification, which is persistent throughout the summer, and is punctuated with occasional coastal upwelling events (Botas et al. 1990). These coastal features develop as a consequence of the prevailing winds off the Cantabrian coast, which are northeasterly during summer and generate westward superficial currents, producing an Ekman transport offshore (Botas et al. 1990). Previous studies in this area have shown changes in phytoplankton ecophysiology (Marañón & Fernández 1995), microplankton trophic structure (Bode & Fernández 1992), and heterotrophic abundance and production (Barquero et al. 1998) as a response to coastal upwelling during summer. However, the bacterial response to a short-term environmental variability, as consequence of sporadic upwelling pulses, still remains to be explored.

The aim of this study was to estimate the role of bacterial activity in carbon cycling in an oligotrophic environment. Although oligotrophic regions are low productivity areas, they cover ~70% of the total marine surface, and may represent up to 40% of total annual production (Iverson et al. 2000). Therefore, the assessment of BGE for natural planktonic bacteria in oligotrophic environments is ecologically important and will also improve our understanding of the biogenic carbon fluxes in marine ecosystems. The purpose of the present work is to provide some accurate growth-efficiency data of bacterioplankton in the open sea. In order to meet this main objective, we report the results of community and bacterial metabolism in 2 contrasting scenarios: before and after an upwelling pulse, during the oligotrophic phase in a temperate sea.

MATERIALS AND METHODS

Sampling. The study was conducted at a shelf-break station in the southern Bay of Biscay (43°42.00' N to 06°42.00' W, 180 m bottom depth) during summer 1999. This site is far from coastal nutrient supplies and

oligotrophic conditions therefore prevail. During the 3 visits to the sampling station, the vertical distribution of photosynthetically available radiation (PAR, 400 to 700 nm) was determined with a submersible spherical quantum sensor (LiCor LI-193SA). Vertical profiles of temperature and conductivity were obtained with a SBE25-03 CTD instrument. Water samples were collected from 4 to 10 discrete depths with 5 l Teflon General Oceanics Niskin bottles.

Samples for the determination of dissolved nitrate (NO_3) concentration were frozen immediately after collection and stored at -20°C until analysis, which was done colourimetrically using a Technicon AAI Autoanalyser. The concentration of nitrate was measured according to the methods described by Grasshoff et al. (1983).

Water samples for dissolved organic carbon (DOC) analysis were filtered through Whatman GF/F glass-fibre filters. The filtrates were collected directly in pre-combusted (450°C for 24 h) glass ampoules with 50 μl H_2SO_4 (85%), sealed immediately after collection and stored in the dark at 4°C until analysis. DOC concentration was determined by high temperature (680°C) oxidation with a Shimadzu 5000A TOC analyzer, as described by Doval et al. (1997).

With the aim of obtaining a time series of upwelling intensity in the study area, Ekman transport values (M) were calculated each day using the equation of Bakun (1973):

$$M = \frac{\tau}{f} = \frac{\rho_a c_d v \sqrt{u^2 + v^2}}{2\Omega \sin\Phi}$$

where ρ_a is the density of air (1.22 kg m^{-3} at 15°C), c_d is a dimensionless, empirical drag coefficient (1.4×10^{-3}), u is the north-south component of wind speed, v is the east-west component of wind speed (this is the component that induces upwelling in the Cantabrian coast), τ is wind stress (N m^{-2}), f is the Coriolis parameter ($9.96 \times 10^{-5} \text{ s}^{-1}$ at 43°N), Ω is angular velocity (rad s^{-1}), and Φ is latitude (degrees). Wind data were recorded at the Asturias airport, which is located near the coast in the vicinity of the sampling station. Positive values of M indicate offshore Ekman transport, which produces upwelling of cold, nutrient-charged, subsurface waters. Negative Ekman transport values ($-M$) indicate downwelling of surface water on the coast, which increases the thermal stratification conditions in the water column. The temporal distribution of surface temperature was measured by a continuous data logger (ONSET Computer Corporation)

Biomass and abundance of phytoplankton and bacterioplankton. Chl *a* concentration was estimated fluorometrically with a Turner Designs 10-AU fluorometer. Samples of 200 ml were filtered onto 0.2 μm polycarbonate filters for the estimation of total chl *a*

concentration. All filters were frozen upon collection and pigments were extracted in 5 ml of 90 % acetone overnight at 4°C (Strickland & Parson 1972).

For the identification and enumeration of microplankton, seawater samples were collected in 125 cm³ glass bottles with Lugol's iodine. Microplankton enumeration and identification was carried out with an inverted microscope following the methodology of Utermöhl (1958). In order to convert phytoplankton (>2 µm) and ciliate abundance to carbon biomass, we used the following equation:

$$\log(\text{pg C cell}^{-1}) = \log a + b \times \log V (\mu\text{m}^3)$$

where $\log a$ is the y -intercept, b is the slope, and V is volume (see Fig. 4 in Menden-Deuer & Lessard 2000). Samples for picoalgal and bacterial abundance were fixed with 1 % paraformaldehyde and 0.05 % (final conc.) of glutaraldehyde and then frozen in liquid N₂. Samples were stored at -70°C until analysis. Picoplankton abundances (heterotrophic bacteria, *Prochlorococcus*, *Synechococcus* and picoeukaryotes) were measured using a bench flow cytometer (Becton & Dickinson FACScalibur) equipped with a laser emitting at 488 nm. Cyanobacteria and picoeukaryotes were identified in plots of side scatter (SSC) versus orange fluorescence (FL2) and red fluorescence (FL3). The standard was run at high speed (~60 ml min⁻¹) in every sample and terminated once 10 000 events had been recorded. The samples for heterotrophic bacteria were stained with Syto-13 Molecular Probes (2.5 mM) and we also added, as an internal standard, an aliquot of fluorescent Latex beads (0.96 µm, Polysciences) previously counted by epifluorescent microscopy. In order to convert cell abundance to carbon biomass we used the following conversion factors: 20 fg C cell⁻¹ for heterotrophic bacteria (Lee & Fuhrman 1987), 2100 fg C cell⁻¹ for picoeukaryotes (Campbell et al. 1994) and 103 fg C cell⁻¹ reported by Zubkov et al. (1998) for *Synechococcus*.

Phytoplankton production and community metabolism. Particulate organic carbon production (POC-prod) rates were determined by means of simulated *in situ* incubations with ¹⁴C. Incubation bottles were cleaned by soaking them in 1N HCl overnight and then rinsing them with deionized water. Seawater samples were dispensed into 70 ml Pyrex bottles inoculated with ~370 kBq (~10 µCi) of NaH¹⁴CO₃ and incubated for 24 h. Three light bottles and 1 dark bottle were incubated for each depth. Temperature and irradiance were simulated taking into account the information obtained previously from the vertical PAR and CTD profiles. Samples were incubated at *in situ* temperature. *In situ* irradiance was simulated using a set of blue plastic filters placed at the top of the incubators. The samples were then filtered under low vacuum

pressure (<100 Hg) through 0.2 µm polycarbonate filters. Filters were exposed to concentrated HCl fumes overnight for degassing and then placed in scintillation vials with 3 ml of scintillation cocktail. Radioactivity in each sample was determined on a Wallac 1409 DSA Liquid Scintillation Counter and quenching was corrected using an internal standard.

For the estimation of the rates of oxygen production and consumption, seawater samples were transferred into calibrated borosilicate glass bottles with a nominal volume of 130 cm³. For each station and depth, 6 bottles were fixed immediately (zero bottles), 6 additional bottles were placed in opaque plastic boxes at the bottom of the incubators (dark bottles), and 6 more were placed at the top of the incubators and received the simulated irradiance (light bottles). After 24 h, the light and dark samples were fixed. The concentration of the dissolved oxygen was determined by Winkler titration using a PC-based system with a photometric endpoint detector (Williams & Jenkinson 1982). Community respiration rate (R ; [O₂]_{zero bottle} - [O₂]_{dark bottle}), NCP ([O₂]_{light bottle} - [O₂]_{zero bottle}), and GPP (= NCP + R) were then calculated as µmol O₂ l⁻¹ d⁻¹. We obtained a pooled variation coefficient of 0.092 % ± 0.008 (mean ± SE, $n = 59$, range = 0.01 - 0.31).

Bacterial metabolism. In order to measure the rates of bacterial respiration, seawater samples were filtered, using a gentle reverse filtration system (0.8 µm, Williams 1981), in order to remove most phytoplankton and protozoa. We did not filter more than 20 l of a given sample through a filter at each depth and we did not find any significant differences in DOC concentration (Student's $t_{11} = 1.61$, $n = 12$, $p = 0.14$), nitrate concentration (Student's $t_7 = 0.61$, $n = 8$, $p = 0.14$), or bacterial abundance (Student's $t_{11} = 0.81$, $n = 12$, $p = 0.43$) between filtered (<0.8 µm) and whole seawater, suggesting that most of the heterotrophic bacterial assemblage passed through the 0.8 µm filter. The percentage of chl *a* measured within the filtered samples was ca. 5 % of the chl *a* measured within the unfiltered samples, which suggests that most of the phytoplankton, including picoautotrophs were retained in the filter. The filtered water was used to measure BR (O₂ consumption), using the same methodology described above for total R .

We estimated BNP within the unfiltered (BNP) and filtered samples (BNP_f) as the rate of ³H-leucine incorporation using standard methods in Eppendorf vials. The replicates from each depth were spiked with 40 nM ³H-leucine and incubated in the dark at *in situ* temperatures (incubator described above) for ~ 4 h. We applied 2 different conversion factors (CFs): the theoretical CF of 3.1 kg C mol leucine⁻¹ (Simon & Azam 1989) and the empirical CF of 0.77 kg C mol⁻¹ (Barbosa et al. 2001) estimated in an ecologically similar area

during late summer. BGE (and BGE_f) is calculated as the ratio of BNP (and BNP_f) to bacterial carbon demand (BCD = [BNP or BNP_f] + BR), assuming a respiration quotient (RQ) of 1 based on the respiration of organic matter having Redfield stoichiometry. The SE for each BGE datum has been calculated from random numbers generated from a Gaussian distribution. Using the mean and SE of each value of BR and BNP, we generated 1000 random numbers from a Gaussian distribution. From each of the 3 BNP and 6 BR replicates, we calculated a BGE datapoint. When we had calculated 100 BGE estimations, we could obtain the SE for each BGE for any depth and month.

We used non-parametric tests as Wilcoxon signed-ranks tests and Spearman coefficients for those cases when data were not normality distributed.

RESULTS

Physical, chemical and biological characteristics during summer

The general vertical and temporal patterns of temperature, nutrients, chlorophyll and total metabolism at the shelf station have been previously described by

Serret et al. (1999). Table 1 summarises the characteristics during the summer of 1999 in terms of physical, chemical and biological parameters, providing the average values for the euphotic layer. The upper mixed-layer temperature increased from June to July and by August had decreased by approximately 1°C. The oxygen saturation in June and July was slightly below 100% except in surface waters, whereas supersaturated waters prevailed at all depths in August. Average DOC concentrations did not show significant differences between the months, but were positively correlated with temperature ($r = 0.74$, $p < 0.01$, $n = 12$). The dissolved nitrate concentration through summer ranged from less than 0.5 μM in surface waters to more than 3 μM in deep waters. We found slightly higher surface nitrate levels in August as compared to June and July. In general, the chl *a* concentration and POC-prod rate did not show a significant difference between the months studied, and were always lower than 0.5 mg m^{-3} and 3 $\text{mg C m}^{-3} \text{d}^{-1}$, respectively. Phytoplankton cell (> 2 mm) numbers were higher in August, with diatoms accounting for ~50% of total phytoplankton cell abundance. Total phytoplankton abundance was also positively correlated with the O₂ concentration ($r = 0.72$, $p < 0.005$, $n = 12$). In contrast, the relative con-

Table 1. Average values (\pm SE) within the photic layer of selected physical, chemical and biological variables at the stations sampled during summer. Variables measured include: temperature of the upper mixed layer (T_{uml}), temperature of the deep waters (T), percentage of oxygen saturation (Sat, %), dissolved organic carbon (DOC), surface nitrate (NO₃ surface), nitrate (NO₃), total chl *a*, particulate organic carbon production (POC-prod), phytoplankton abundance (PA), relative contribution of diatoms (%) and dinoflagellates (%) to phytoplankton abundance (> 2 mm), heterotrophic bacteria abundance (BA), gross primary production (GPP), net community production (NCP), community respiration (R), bacterial respiration (BR), total bacterial net production (BNP), filtered bacterial net production (BNP_f), and bacterial growth efficiencies (BGE, BGE_f [filtered])

	June	July	August
T_{uml} (°C)	15.6	19.7	18.6
T (°C)	12.1	12.2	13.3
Sat (%)	97.7 \pm 2.2	99.8 \pm 2.5	105.8 \pm 2.5
DOC (μM)	85.0 \pm 5.3	90.5 \pm 6.5	84.0 \pm 8.5
NO ₃ surface (μM)	0.1	<0.1	0.3
NO ₃ (μM)	0.6 \pm 0.5	3.6 \pm 1.4	2.9 \pm 1.0
Chl <i>a</i> (mg m^{-3})	0.4 \pm 0.1	0.2 \pm 0.1	0.5 \pm 0.2
POC-prod ($\text{mg C m}^{-3} \text{d}^{-1}$)	2.7 \pm 1.9	0.9 \pm 0.5	1.0 \pm 0.6
PA (> 2 mm, cells ml ⁻¹)	21 \pm 7	29 \pm 11	40 \pm 11
PA (< 2 mm, $\times 10^3$ cells ml ⁻¹)	27.9 \pm 16.4	5.4 \pm 2.4	18.5 \pm 10.7
Diatoms (%)	15 \pm 1	14 \pm 7	51 \pm 7
Dinoflagellates (%)	71 \pm 3	75 \pm 7	43 \pm 7
Ciliates (cells ml ⁻¹)	2.1 \pm 0.9	4.1 \pm 2.2	2.9 \pm 0.9
BA ($\times 10^6$ cells ml ⁻¹)	0.8 \pm 0.2	0.7 \pm 0.2	0.5 \pm 0.1
GPP ($\text{mmol O}_2 \text{m}^{-3} \text{d}^{-1}$)	0.3 \pm 0.1	0.2 \pm 0.1	1.1 \pm 0.4
NCP ($\text{mmol O}_2 \text{m}^{-3} \text{d}^{-1}$)	-1.6 \pm 0.8	-2.2 \pm 1.0	-1.3 \pm 1.1
R ($\text{mmol O}_2 \text{m}^{-3} \text{d}^{-1}$)	1.9 \pm 0.8	2.4 \pm 0.92	2.5 \pm 1.4
BR ($\text{mmol O}_2 \text{m}^{-3} \text{d}^{-1}$)	1.5 \pm 0.5	2.4 \pm 1.0	1.4 \pm 1.3
BNP (pmol leucine l ⁻¹ h ⁻¹)	15.3 \pm 1.6	11.8 \pm 2.7	86.7 \pm 76.9
BNP _f (pmol leucine l ⁻¹ h ⁻¹)	7.9 \pm 3.0	6.9 \pm 2.3	16.9 \pm 14.5
BGE (%) (BGE _f %)	2.7 \pm 1.3 (0.9 \pm 0.2)	1.0 \pm 0.4 (0.7 \pm 0.2)	8.0 \pm 0.6 (1.9 \pm 1.0)
BGE _f (%)	0.9 \pm 0.2 (3.4 \pm 0.7)	0.7 \pm 0.3 (2.7 \pm 0.9)	1.9 \pm 1.0 (7.1 \pm 3.6)

tribution of dinoflagellates to total phytoplankton abundance during June and July was significantly higher (>70%) than the diatom contribution (<20%). Ciliate abundance was relatively low and was positively correlated with the abundance of dinoflagellates ($r = 0.82$, $p < 0.001$, $n = 12$). The number of picoautotrophs was generally above 10^3 cells ml^{-1} , whereas heterotrophic bacterial abundance (BA) was generally above 10^5 cells ml^{-1} , which decreased as summer progressed. BA was positively correlated with temperature ($r = 0.59$, $p < 0.05$, $n = 12$) and DOC concentration ($r = 0.60$, $p < 0.05$, $n = 12$), and was negatively correlated with nitrate concentration ($r = -0.78$, $p < 0.005$, $n = 12$).

Community and bacterial metabolism

GPP rates were higher in August (Table 1) than in June or July. GPP was correlated with diatom abundance ($r = 0.69$, $p < 0.05$, $n = 12$) and was significantly lower (mean \pm S.E., 0.54 ± 0.17) than R (2.26 ± 0.56 $\text{mmol O}_2 \text{ m}^{-3} \text{ d}^{-1}$, $F_{1,22} = 16.24$, $p < 0.005$), implying that the community metabolic balance was net heterotrophic during the stratification period (NCP < 0). Pre-filtration of seawater samples through $0.8 \mu\text{m}$ filters allowed us to assess the metabolism of the bacterial community. We found that respiration rates in filtered samples (BR) were significantly lower than in the unfiltered samples (R ; Student's $t_{11} = 2.43$, $n = 12$, $p < 0.05$). BR showed a highly significant, positive correlation with R and with DOC concentration (Fig. 1A, $r^2 = 0.80$; Fig. 1B, $r^2 = 0.87$). On average, BR

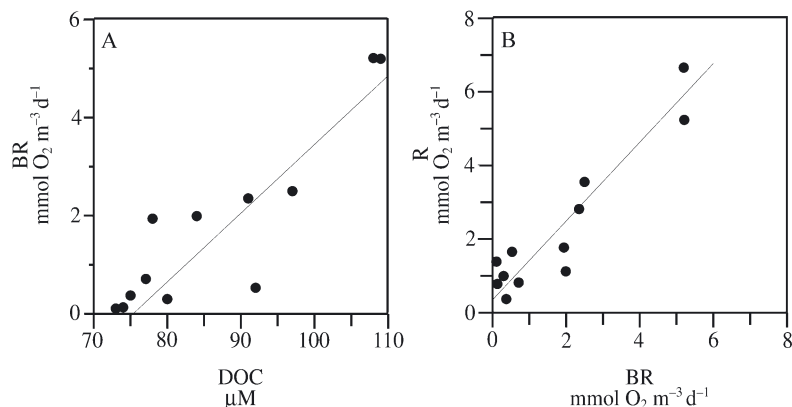


Fig. 1. Parameters of linear regression model, of the form $y = a(x) + b$. (A) Bacterial respiration (BR) and dissolved organic carbon concentration (DOC). The fitted line is the reduced major axis (rma) regression. Values are: x (DOC), y (BR), a (0.14 ± 0.09), b (-10.55 ± 3.89), $r^2 = 0.80$, $n = 12$, $p < 0.0001$. (B) Community respiration (R) and BR. The fitted line is the rma regression. Values are: x (BR), y (R), a (1.07 ± 0.06), b (0.35 ± 0.66), $r^2 = 0.87$, $n = 12$, $p < 0.0001$

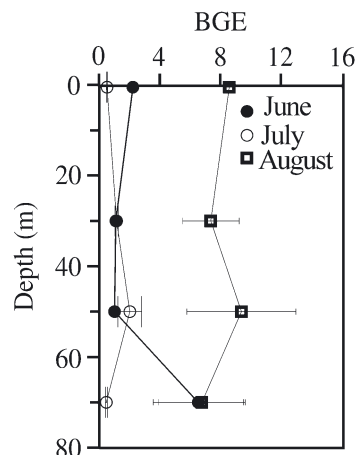


Fig. 2. Vertical distribution of bacterial growth efficiencies (BGE; mean \pm SE) during, June, July and August. The coefficient factor used is $\text{CF} = 0.77 \text{ kg C mol}^{-1}$

accounted for 74% ($\pm 14\%$) of R . BNP_f was significantly lower than BNP (Wilcoxon signed ranks test, $Z = 3.94$, $n = 12$, $p < 0.003$). Bacterial biomass and BNP showed a positive correlation in unfiltered and filtered samples (Spearman coefficient, $r_s = 0.61$, $p < 0.05$, $n = 12$; $r_s = 0.71$, $p = 0.01$, $n = 12$, respectively). BR (assuming a RQ of 1) was significantly higher than BNP (assuming a CF of 0.77, $F_{1,21} = 10.56$, $p < 0.005$, $n = 22$; and $\text{CF} = 3.1$, $F_{1,21} = 9.49$, $p < 0.01$, $n = 22$), which yielded BGEs of lower than 10 and 30%, respectively. BGE was significantly higher than BGE_f (using the $\text{CF} = 0.77$, Student's $t_{11} = 2.88$, $n = 12$, $p = 0.01$). BGE was positively correlated with GPP ($r = 0.69$, $p < 0.05$) and showed a temporal variability, where BGE was significantly higher in August than in June and July ($F_{1,9} = 18.34$, $p < 0.005$, $n = 12$; Fig. 2). However, BGE_f was not correlated with GPP and did not show clear differences between months.

In order to explore the causes of variability in the physical, chemical and biological characteristics during summer, we plotted a time series of upwelling intensity and temporal distribution of surface temperature in the study area. Temporal distribution of surface temperature (Fig. 3) showed a marked decrease of temperature and high positive Ekman transport values at the end of July, immediately before the sampling in August. Another high positive Ekman transport value was measured at the end of June, which occurred immediately before the sampling in July. However, this upwelled water did not reach the surface, evident because there was no important decrease in the surface water temperature.

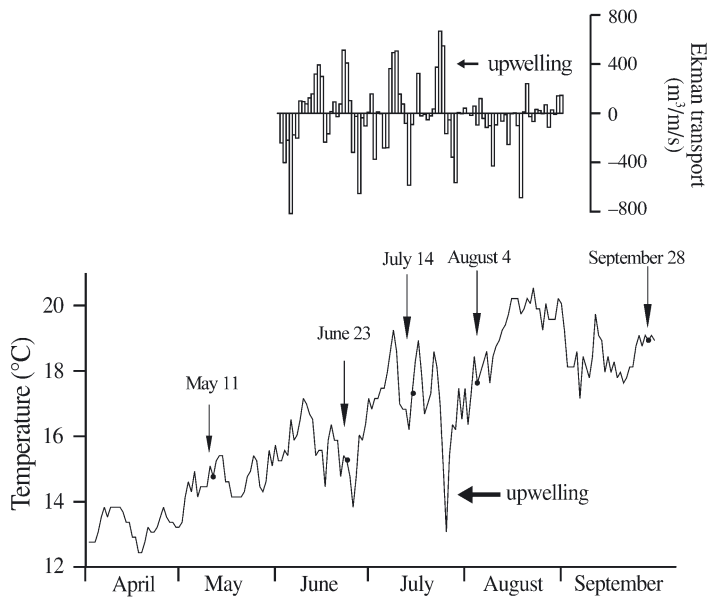


Fig. 3. Temporal distribution of surface temperature and Ekman transport in the coastal region of the Southern Bay of Biscay (Campiello)

Carbon flux through the microbial community: before and after an upwelling event

In order to better appreciate differences in community metabolism during the summer, we have considered the euphotic depth-integrated estimates in 2 contrasting scenarios: pre- and post-upwelling (Table 2). Although heterotrophy dominated during the summer, in the post-upwelling phase the magnitude of imbalance between R and GPP was lower due to a higher GPP. The upwelling pulse may also have had an effect on bacterial community (Table 2), decreasing bacterial abundance and respiration but also increasing BNP (and BNP_f). The relative contribution of bacteria to R decreased after the upwelling pulse. On the other hand, the bacterial biomass to phytoplankton biomass ratio was lower after the upwelling event, when GPP was significantly higher (Table 2). The bacterial production to phytoplankton production ratio showed that BNP averaged 11% (44% using $CF = 3.1 \text{ kg C mol}^{-1}$) of primary production during summer. We have estimated this ratio using the GPP instead of the POC-prod value because our incubations lasted 24 h, so POC-prod will lie closer to net than gross production. This ratio also showed temporal variability and a contrasting trend to that of the biomass ratio, as the bacteria to phytoplankton production ratio increased as GPP increased (Table 2).

DISCUSSION

Whole and filtered samples

The pre-filtration method to estimate BR has been widely used (e.g. Biddanda et al. 1994, Blight et al. 1995, Smith & Kemp 2001); however, the possibility of artefacts interfering with the results is always a concern. Long-term bottle incubations (more than 24 h) measure the activity of opportunistic assemblages and not the original bacterial community (Massana et al. 2001) and also can change the bacterial physiology (Malone et al. 1991). In addition, removing the bacterial predators can overestimate the microbial activity (e.g. Pomeroy et al. 1994). However, our incubation lasted 24 h, which is the minimum time required to detect changes in oxygen concentration under oligotrophic conditions. According to Massana et al. (2001) this incubation length should result in minimal changes in the assemblage. Respiration and bacterial activity were also always significantly lower within the filtered samples, suggesting that the problems of overestimation were likely relatively minor.

The response of the bacterial assemblage to the filtration treatment after 4 h of incubation is, on the other hand, interesting in view of the fact that bacterial production was significantly higher within the whole seawater samples than in the filtered samples. These results help us to understand whether autochthonous labile DOC production is an important carbon source to maintain bacterial activity. Labile DOC production originates from several sources, such as phytoplankton exudation (e.g. Fogg 1983, Nagata 2000), grazing phytoplankton (Jumars et al. 1989, Strom et al. 1997), zooplankton and protist release (Lee & Fisher 1994),

Table 2. Depth-integrated estimates of gross primary production ($\int GPP$), net community production ($\int NCP$), community respiration ($\int R$), bacterial abundance ($\int BA$), bacterial respiration ($\int BR$), filtered bacterial net production ($\int BNP_f$), total bacterial net production ($\int BNP$), bacteria to phytoplankton biomass ratio ($\int BB:\int PB$), bacteria net- to phytoplankton gross-production ($\int BNP:\int GPP$), at the 2 contrasting scenarios pre- and post-upwelling during summer. Values in parentheses are rates calculated with conversion factors = $3.1 \text{ kg C mol}^{-1}$

	Pre-upwelling	Post-upwelling
$\int GPP$ ($\text{mg C m}^{-2} \text{ d}^{-1}$)	234	904
$\int NCP$ ($\text{mg C m}^{-2} \text{ d}^{-1}$)	-1802	-1189
$\int R$ ($\text{mg C m}^{-2} \text{ d}^{-1}$)	2036	2093
$\int BA$ ($\times 10^9 \text{ cells m}^{-2}$)	54	38
$\int BR$ ($\text{mg C m}^{-2} \text{ d}^{-1}$)	1743	1213
$\int BNP_f$ ($\text{mg C m}^{-2} \text{ d}^{-1}$)	11 (45)	18 (74)
$\int BB/\int PB$	0.98	0.42
$\int BCD/\int GPP$	7.06 (6.8)	1.7 (1.3)

dissolution of fecal pellets (Lampitt et al. 1990), cell lysis by viruses (Fuhrman 1992, Bratbak et al. 1998), autolysis (Agustí et al. 1998) photochemical degradation of recalcitrant DOC (Tranvik 1992), and DOC production by heterotrophic bacteria (Iturriaga & Zsolnay 1981). Filtration removes the substrates and nutrient sources produced by organisms $> 0.8 \mu\text{m}$, reduces the grazing impact and might increase the release of DOC as a result of damage to plankton. However, differences in the initial DOC concentration between filtered and unfiltered samples were not significant, suggesting that any DOC contamination during the filtration process is negligible. It is possible that some of the most active bacteria were retained in the filter because they were attached to particles. According to Gasol et al. (1999), bacteria with a high DNA content (HDNA) compared to those with a low DNA content (LDNA) represent the active bacteria in plankton. However, we did not see significant differences in the HDNA-bacteria concentration between filtered and unfiltered samples (data not shown, Student's $t_{11} = 0.23$, $n = 12$, $p = 0.82$). Also, counting by use of flow cytometry could underestimate bacterial abundance when many bacteria are attached. However, bacteria attached to particles have been considered to make up less than 10% of the total bacterial community (Griffith et al. 1994) and Lemée et al. (2002) did not find bacteria attached to particles during summer. According to several authors (e.g. Sanders et al. 1992, Vaqué et al. 1994, Caron et al. 2000), bacterivorous grazers are the main loss-factor for bacteria. However, the higher bacterial production within the unfiltered samples indicated that grazing pressure on bacteria was probably low, otherwise bacterial production should have been higher within the filtered samples. The difference in bacterial production between filtered and unfiltered samples was $\sim 40\%$ before the upwelling and $\sim 80\%$ after the upwelling. This percentage of bacterial production must have been maintained by the DOC release caused by microzooplankton grazing on phytoplankton, plus phytoplankton exudation and cell lysis (viral infection and autolysis). However, according to Anderson & Ducklow (2001), the latter supply route of DOC is a major process that fuels bacterial production.

Temporal variability of bacterial metabolism

During the stratified oligotrophic period within temperate waters, wind-mixing events may temporarily enhance the vertical flux of nutrients across the pycnocline to the surface layer (Eppley & Renger 1988), thus causing a transient shift in the pelagic food-web structure. According to Legendre & Rassoulzadegan (1996), such a temporal discontinuity will give rise to an

increase in the flow of biogenic carbon-export relative to primary production, decreasing the relative contribution of picoplankton-mediated recycling processes. The Bay of Biscay displays the characteristics of an oligotrophic system during summer, when nutrient concentrations in surface waters are low and material discharge from rivers decreases (Prego & Vergara 1998). However, this period of thermal stratification is disrupted by periodic coastal upwelling events. The effect of these upwelling episodes is difficult to model, due to their changing intensity, sporadic occurrence and variable biological response. The present short-term environmental changes during summer allowed us to assess the effect of upwelling on bacterial metabolism. Bacterial abundance measured in this study (0.5 to 0.8×10^6 cell ml^{-1}) was comparable with the mean value for oligotrophic systems (0.5 to 1×10^6 cell ml^{-1}) presented in the recent review by Cotner & Bidanda (2002). On the other hand, BNP computed for oligotrophic regions was in the range 0.010 to $11 \text{ mg C l}^{-1} \text{ d}^{-1}$ (Ducklow & Carlson 1992), which includes our values (0.11 to $1.33 \text{ mg C l}^{-1} \text{ d}^{-1}$, assuming both CFs). The positive correlations between BA and both temperature and DOC, and the negative correlation with nutrients, is in agreement with a reduction in bacterial abundance after the upwelling event, when temperature and DOC slightly decreased and nutrients increased. The relative contribution of heterotrophic bacteria biomass to total phytoplankton biomass (Table 2) decreased as the productivity of the system increased (post-upwelling). Such a scenario is in agreement with earlier observations (e.g. Cole et al. 1988, del Giorgio et al. 1997). However, the ratio of bacterial production to phytoplankton production showed the opposite trend (Table 2), increasing during the relaxation of the upwelling pulse. These results suggest that after the upwelling event, bacterial biomass could be controlled by grazers (Sanders et al. 1992) and/or viral infection (Weinbauer & Peduzzi 1995, Gasol & Duarte 2000). Our previous results showed that grazing pressure on bacteria was not important, thus viral pressure, which generally increases after a phytoplankton bloom (e.g. Bratbak et al. 1995), should be considered as a factor controlling bacterial biomass.

Mean R ($\sim 2 \text{ g C m}^{-2} \text{ d}^{-1}$) was close to the average reported in a recent revision ($1.2 \text{ g C m}^{-2} \text{ d}^{-1}$, assuming a RQ of 1) for oligotrophic environments (del Giorgio & Duarte 2002). During summer, BR averaged 70% of R and was strongly correlated with DOC concentration. These results suggest that bacteria are the major contributors to bulk R , and that variability in BR is largely attributable to DOC concentration and plays an important role in organic matter mineralization. BR was not constant and its contribution to R declined after the

upwelling pulse, whereas phytoplankton biomass and GPP increased in response to the nutrient addition to the surface waters.

Carbon flux through bacterioplankton under oligotrophic conditions

The results obtained in the present study have allowed us to build the biogenic carbon budget through the microbial community during summer before and after the upwelling event (Fig. 4). This diagram shows a shift of the microplankton community in

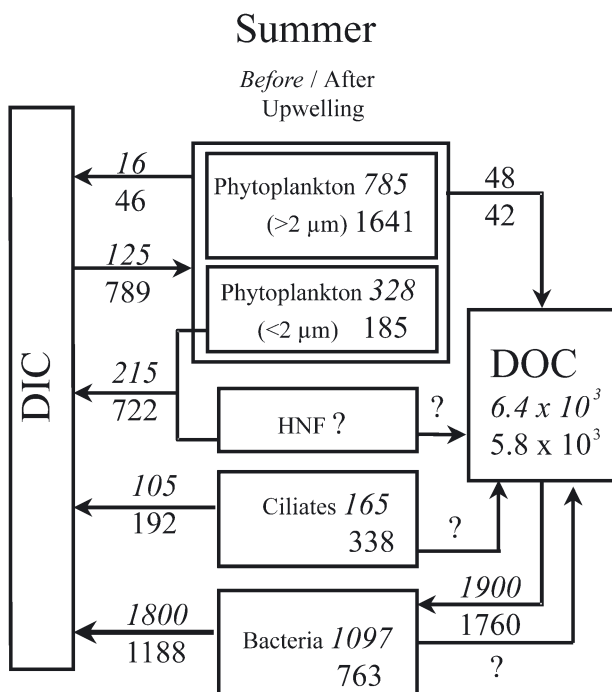


Fig. 4. Carbon flux through the microbial community estimated for pre (italics) and post- (non-italics) upwelling conditions during summer. Boxes indicate the biomass of the different microplankton groups, expressed in mg C m⁻². Arrows represent photic depth-integrated carbon fluxes, expressed in mg C m⁻² d⁻¹. Daily light period was calculated using the equation given by Straskraba & Gnauck (1985). Oxygen fluxes were transformed to carbon fluxes using a respiration quotient (RQ) of 1. Phytoplankton (>2 μm) respiration (R) was calculated through the allometric equation: R (pl O₂ cell⁻¹ h⁻¹) = 0.0068 × V (μm⁻³)^{0.88} described by Tang & Peters (1995). Ciliate respiration was obtained using the allometric equation, $\log_{10} R$ (nl O₂ cell⁻¹ h⁻¹) = $\log_{10} \text{vol}$ (μm³) × 0.75 – 4.09, proposed by Fenchel & Finlay (1983). Picoalgae and heterotrophic nano-flagellate (HNF) respiration was obtained as the microbial community respiration minus the respiration of ciliate, heterotrophic bacteria and nanoalgae. To calculate the phytoplankton dissolved organic carbon (DOC) production rate we used the following equation (POC-prod: particulate organic carbon production): $\log_{10} \text{DOC}$ (mg C m⁻³ h⁻¹) = $g_{10} \text{POC-prod}$ (mg C m⁻³ h⁻¹) × 0.52 – 0.74, proposed by Teira et al. (2001)

response to an upwelling pulse. Phytoplankton in summer was dominated by cells smaller than 2 μm (picoautotrophs, Table 1). However, larger phytoplankton biomass dominates the total phytoplankton carbon pool during summer, and its contribution to the total phytoplankton biomass increased after the upwelling pulse from 70 to 90% (Fig. 4). DOC concentration decreased after the upwelling pulse, whereas the microzooplankton biomass (ciliates) increased. Bacterial production was always significantly lower than BR, giving rise to a BGE lower than 10%, considering 0.77 kg C mol leucine⁻¹ as the coefficient factor more appropriated for this study. The significance of this low BGE is that a high quantity of carbon passes through the bacterial community and is mainly respired, thus not available to higher trophic levels (Legendre & Rassoulzadegan 1995). Before the upwelling event, when conditions were more oligotrophic, most of the carbon consumed by bacteria was respired as CO₂, whereas after the upwelling a higher percentage of this DOC was converted to POC. According to del Giorgio & Cole (1998), the degree of uncoupling between bacterial production and respiration explains the variability observed in BGE during summer. BGE was significantly higher than BGE_r, which indicates that the efficiency at which bacteria consume DOC increases in the presence of DOC producers, such as grazers, virus and phytoplankton. Given that BGE was positively correlated with GPP and was also significantly lower before (2.7 ± 1.3%) than after (8.0 ± 0.6%) the upwelling, this indicates that BGE responds to changes in phytoplankton production, increasing as the system becomes more productive. The mean BR rate per bacteria after the upwelling (0.39 ± 0.07 fmol O₂ h⁻¹ cell⁻¹, except at the surface where BR rate per bacteria was 6.5 fmol O₂ h⁻¹ cell⁻¹, data not shown) was comparable with the mean value (0.1 ± 0.3 fmol O₂ h⁻¹ cell⁻¹) proposed by Biddanda et al. (1994) and was lower than that before the upwelling (3.0 ± 0.89 fmol O₂ h⁻¹ cell⁻¹). However, after the upwelling, when nutrient availability and perhaps the quality of DOC increased, the energetic cost of bacterial growth appeared to decrease and BGE increased. BCD on average always exceeded GPP during summer; however, the ratio of BCD to GPP decreased after the upwelling pulse. These results indicate that phytoplankton sustained most of the BCD after the upwelling pulse, whereas during the period of oligotrophy, additional organic carbon sources maintained bacterial metabolism.

Acknowledgements. We are deeply indebted to E. Maraño for his constructive criticism, helpful suggestions on the manuscript and general encouragement. Thanks are given to J. M. Gasol for helping with the flow cytometry facility and for

his constructive criticism. We also thank X. A. Álvarez-Salgado for measuring DOC samples and J. Sostres and E. Cabal for their technical support, as well as the crew of RV 'José Rioja' for excellent support. I also acknowledge the comments of anonymous referees on a previous draft. This work was supported by CANIGO Grant CE-96-MAS3-CT-0060 and is part of a thesis from the University of Oviedo.

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Editorial responsibility: Fereidoun Rassoulzadegan, Villefranche-sur-Mer, France

*Submitted: January 28, 2003; Accepted: June 30, 2003
Proofs received from author(s): August 21, 2003*